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Ozone and SOA formation potential based on photochemical loss of VOCs during the Beijing summer $\stackrel{\text{}}{\sim}$

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ABSTRACT

Volatile organic compounds (VOCs) are easily degraded by oxidants during atmospheric transport. Therefore, the contribution of VOCs to ozone (O_3) and secondary organic aerosol (SOA) formation at a receptor site is different from that in a source area. In this study, hourly concentrations of VOCs and other pollutants, such as O₃, NOx, HONO, CO, and PM_{2.5}, were measured in the suburbs (Daxing district) of Beijing in August 2019. The photochemical initial concentrations (PICs), in which the photochemical losses of VOCs were accounted for, were calculated to evaluate the contribution of the VOCs to O_3 and SOA formation. The mean (±standard deviation) measured VOC concentrations and the PICs were 11.2 \pm 5.7 and 14.6 \pm 8.4 ppbv, respectively, which correspond to O_3 formation potentials (OFP) of 57.8 \pm 26.3 and 103.9 \pm 109.4 ppbv and SOA formation potentials (SOAP) of 8.4 ± 4.1 and $10.3 \pm 7.4 \ \mu g \ m^{-3}$, respectively. Alkenes contributed 80.5% of the consumed VOCs, followed by aromatics (13.3%) and alkanes (6.2%). The contributions of the alkenes and aromatics to the OFP_{PICs} were 56.8% and 30.3%, respectively; while their corresponding contributions to the SOAP_{PICs} were 1.9% and 97.3%, respectively. The OFP_{PICs} was linearly correlated with the observed O₃ concentrations (OFP_{PICs} = $41.5 + 1.40 \times$ $c_{0,3}$, $R^2 = 0.87$). The O₃ formation was associated with a VOC-limited regime at the receptor site based on the measured VOCs and changed to a transition regime and a NOx sensitive regime based on the PIC. Our results suggest that more attention should be paid to biogenic VOCs when studying O₃ formation in summer in Beijing, while the control of anthropogenic aromatic compounds should be given priority in terms of SOA formation.

Carter, 2010).

leading to the net accumulation of O_3 (Chameides et al., 1992). VOCs are precursors of both SOAs and O_3 in the atmosphere (Odum et al., 1997;

The O₃ formation potential (OFP) and secondary organic aerosol

formation potential (SOAP) are widely used to evaluate the relative

contributions of VOC species and sources to ambient O3 and SOA for-

mation, respectively (Alvim et al., 2018; Kumar et al., 2018; Mozaffar

et al., 2020). In practice, the measured VOC concentrations are directly

used to calculate the OFP or SOAP (Zheng et al., 2021). However, the

VOC concentration and composition observed at a receptor site are

likely to be different from those observed in the source region due to the

photochemical degradation of VOCs during transport (Warneke et al.,

1. Introduction

Atmospheric pollution including particulate matter (PM) and surface ozone (O₃), has received intense attention worldwide, due mainly to their adverse effects on human health (Streets et al., 2007; Krizan et al., 2011). As an important contributor to PM, secondary organic aerosols (SOAs) are formed through the photochemical oxidation of volatile organic compounds (VOCs) and condensation of subsequent organic compounds of low volatility (Han et al., 2017). Tropospheric O₃ is produced via complex photochemistry between VOCs and nitrogen oxides (NOx), during which alkyl peroxide radicals (RO₂) and hydrogen peroxide radicals (HO₂) efficiently convert NO to NO₂, subsequently

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Received 20 February 2021; Received in revised form 17 May 2021; Accepted 20 May 2021 Available online 31 May 2021 0269-7491/© 2021 Elsevier Ltd. All rights reserved. 2004). Both O_3 and SOA formation are related to the consumed VOCs, which react with oxidants, such as OH, NO₃, and O₃. OH is one of the essential oxidants and dominates the degradation of VOCs in the atmosphere, subsequently playing an important role in O_3 (Tan et al., 2018) and SOA formation (Wu et al., 2020). Makar et al. (1999) found that the isoprene concentration was underestimated in models by 40% when the degradation of isoprene by OH radicals was ignored in a one-dimensional canopy model. Gao et al. (2021) found that the emission ratios of highly reactive alkenes were larger than those previously reported in industrial zones when the photochemical loss was considered. Therefore, it is necessary to investigate the differences between the source concentration and the receptor site concentration of VOCs, and its effect on the OFP and SOAP.

The concept of source concentration is similar to that of the photochemical initial concentration (PIC) proposed by Shao et al. (2011), which is based on the photochemical-age approach and has been applied to evaluate the effect of photochemical processing on measured VOC levels. If the emissions of primary VOCs are assumed to be constant, and the influence of fresh emissions during transport is neglected (Roberts et al., 1984; Bertman et al., 1995), the photochemical loss of VOCs during atmospheric transport can be calculated using a ratio method. This method calculates the evolution of the ratio between two VOCs with differences in k_{OH} and has been successfully applied to evaluate the performance of stringent air-quality control measures imposed during the Beijing Olympics (Shao et al., 2011). Shao et al. (2011) found that reduced emissions from gasoline vehicles explained 59-68% of the reduction in the OFPs and suggested that the consumption-based OFP was better than the measurement-based OFP for identifying the key reactive species in local O_3 formation. Yuan et al. (2013) used the photochemical-age-based method to calculate VOC emission ratios in Changdao, and the calculated results were in good agreement with the emission inventory. Han et al. (2017) used this method to estimate the SOA production ratio and compared the characteristics of the VOCs and the formation of SOAs between haze and haze free days in Shanghai. However, this method is rarely used to study the impact of PICs on O₃ and SOA formation in suburban areas because O_3 pollution usually occurs in summer while PM pollution occurs in winter (Ma et al., 2021; Wu et al., 2021). Recently, O_3 and PM pollution occurred simultaneously in summer in Beijing (Duan et al., 2020; Li et al., 2021), highlighting the need to synergistically control O_3 and PM pollution in China.

 O_3 concentrations in urban Beijing have exhibited a significant upward trend from May to September, over the period 2001–2015 (Tang et al., 2009; Zhang et al., 2015; Cheng et al., 2018). Moreover, the contribution of the SOA fraction to PM_{2.5} increased by around 1.5 times in the summer of 2015 compared with that in 2008 (Duan et al., 2020). As a result, it is necessary to identify the crucial VOC species that govern the summertime O_3 and SOA formation in Beijing. In this study, we calculated the PIC and assessed the consumption-based O_3 and SOA formation potential in summer in Beijing. The differences between the consumption-based OFP and SOAP and the observation-based values were investigated. The results of this study improve our understanding of the complex photochemistry and provide a scientific basis for efficiently and simultaneously reducing both PM and O_3 pollution in Beijing.

2. Material and methods

2.1. Sampling site description

The sampling site (39.73°N, 116.33°E) is located on the Qingyuan campus of the Beijing Institute of Petrochemical Technology (BIPT) (Fig. 1) and has been described in our previous work (Chen et al., 2020). Briefly, the instruments were installed in a temperature-controlled room on the top floor of a nine-story building (about 27 m above the ground), approximately 600 m from two adjacent streets (Xinghua Street in north-south direction, and Qingyuan Street in east-west direction). The sampling site is a representative suburban site (Chen et al., 2020). Ambient air was drawn through a Teflon manifold (BMET-S, Beijing Saak-Mar Environmental Instrument Ltd.) from the roof using a bypass pump for gas phase pollutant measurements; while particulate matter



Fig. 1. (A) Map of Beijing, (B) Sampling site and (C) a photograph of the laboratory.

was sampled through a $PM_{2.5}$ inlet with polished steel tubing using a bypass pump (16.7 L min⁻¹).

2.2. Instrumentation

Due to the active photochemistry and low rainfall frequency in August, the observations were carried out from August 1 to 28, 2019. The concentrations of non-methane hydrocarbons (NMHCs) were continuously measured using a custom-built gas chromatography flame ionization detector (GC-FID) system, with a time resolution of 1 h. The details of the analytical procedure have been described by Liu et al. (2016). Briefly, the GC-FID system consisted of a cooling unit, a sampling unit, a separation unit, and an FID, operated at 250 °C. A 400 mL sample of air was collected using a CarbopackTM B sorbent (60/80 mesh) filled stainless-steel tube at a flow rate of 40 mL min⁻¹. Then, the adsorption tube was quickly heated to 180 °C for desorbing and injecting the NMHCs into a single column (OV-1, 30 mL \times 0.32 mm I. D.) for separation. The temperature of the capillary column (starting at 60 °C) was held at 60 °C for 3 min, increased to 20 °C at the rate of $12 \circ C \min^{-1}$, increased to 30 °C at the rate of 6 °C min⁻¹, increased to 170 °C at the rate of 10 °C min⁻¹, and then held for 2 min. The C2–C12 NMHCs were quantified using external standard curves derived for 57 NMHCs in gas mixtures balanced with N₂ (Spectra Gases Inc., USA) in the range of 1.0–30.0 ppbv. The R^2 values of the calibration curves were 0.99 for all of the NMHCs, and the GC-FID was calibrated weekly. The variations in the responses of the target species were within 6% of the calibration curve. The method detection limits (MDLs) for the NMHCs ranged from 0.02 to 0.10 ppbv, with relative standard deviations (RSD) of 0.3-4.4% (Liu et al., 2016).

NOx, SO₂, CO, and O₃ analyzers (42i, 43i, 48i, and 49i, respectively, Thermo ScientificTM, USA) were used to measure the concentrations of the corresponding trace gases. A custom-built long path absorption photometer (LOPAP) was used to measure the HONO concentration (Liu et al., 2020c). Briefly, the HONO in ambient air was absorbed using a solution containing 0.06 M sulfnilamide and 1 M HCl, to which *N*-(1-naphthyl) ethylene-diaminedihydrochloride (0.8 mml L⁻¹) was added to generate an azo dye. The azo dye was detected in a Teflon absorption cell (Liquid Core Waveguide, Biogeneral, USA) using a mini-spectrometer with a diode array detector (SD2000, Ocean Optics, USA). The HONO concentration was obtained by subtracting the calibrated signal of the second coil from the first coil and calibrated using external standard nitrile solutions.

The PM_{2.5} mass concentration was measured using a tapered element oscillating microbalance (TEOM-1405DF, Thermo Scientific™, USA). A High-Resolution Time-of-Flight Aerosol Mass spectrometer (HR-TOF-AMS, Aerodyne, USA), which was alternately operated in V-mode and W-mode, was used to measure the composition of the PM_{1.0}. The details have been described in our previous work (Chen et al., 2020). Briefly, the aerosol particles were focused into a tight beam through an aerodynamic lens and were thermally vaporized at 600 °C. The vapors were ionized via electron impact (EI, 70 eV) and were detected using a TOF-AMS detector. The data were analyzed using the PIKA 1.16I software. A collection efficiency (CE) factor of 0.5 and relative ionization efficiency (RIE) values of 1.1, 1.2, 1.4, and 1.3 for NO₃⁻, SO₄²⁻, organic aerosol (OA), and Cl⁻, respectively, were used (Chen et al., 2020). A weather station (AWS310, Vaisala, Finland) was used to measure the meteorological variables, including the temperature (T), pressure (P), relative humidity (RH), wind speed (WS), and wind direction (WD). A J_{NO2} filter-radiometer (J_{NO2} radiometer, Metcon, Germany) was used to measure the photolysis rate (J_{NO2}) by continuously measuring the actinic flux in the wavelength range of 285-375 nm.

2.3. Calculation of photochemical loss of VOCs

The photochemical loss of VOCs was calculated using the ratio method (Shao et al., 2011; Yuan et al., 2013). Changes in VOC

concentrations as a function of time due to photochemical reaction are described in Eq. (1), from which the OH exposure can be calculated as:

$$\int_{0}^{T} c_{OH} dt = \frac{1}{k_{A,OH} - k_{B,OH}} \left[ln \left(\frac{c_A}{c_B} \right)_0 - ln \left(\frac{c_A}{c_B} \right)_t \right]$$
(1)

where $(c_A/c_B)_0$ and $(c_A/c_B)_t$ are the emission ratio and the measured ratio of compound A to compound B, respectively; $k_{A,OH}$ and $k_{B,OH}$ are the second-order reaction rates of compound A to compound B, respectively. Thus, for any VOC_i, the PIC is calculated using Eq. (2).

$$c_{i,0} = \frac{c_{i,t}}{exp(-k_{i,OH})exp\left(\frac{1}{k_{A,OH}-k_{B,OH}}\left[ln\left(\frac{c_A}{c_B}\right)_0 - ln\left(\frac{c_A}{c_B}\right)_t\right]\right)}$$
(2)

and the photochemical loss of VOC is calculated using Eq. (3).

$$\Delta c_{i} = c_{i,t} \left[\frac{1}{exp(-k_{i,OH})exp\left(\frac{1}{k_{A,OH}-k_{B,OH}}\left[ln\left(\frac{c_{A}}{c_{B}}\right)_{0}-ln\left(\frac{c_{A}}{c_{B}}\right)_{t}\right]\right)} - 1 \right]$$
(3)

In this study, the ethylbenzene/xylene pair was used to calculate ambient OH exposure (Shao et al., 2011; Zheng et al., 2021). Justification for the use of ethylbenzene and xylene is as follows. Firstly, the concentrations of xylene and ethylbenzene are well correlated, which indicates that they are simultaneously emitted (Fig. S1); Secondly, the $k_{\rm OH}$ values of ethylbenzene and xylene are 7×10^{-12} and 1.87×10^{-11} cm^3 molecule⁻¹ s⁻¹, respectively (Atkinson and Arey, 2003), which indicates that they have different degradation rates; Finally, the PICs calculated using the ethylbenzene/xylene pair are in good agreement with those calculated using other tracers, such as *i*-butene/propene (Fig. S2). The hourly concentrations of ethylbenzene and xylene were used to calculate the photochemical loss of VOCs. Then, the PIC was calculated by adding the measured VOC concentration and the calculated photochemical loss. In addition to OH exposure, atmospheric transport may also affect the VOC ratio. Sensitivity tests showed that the uncertainty caused by the OH exposure (from -25% to +25%) ranged from 0.77 to 1.33 (Table S1).

2.4. Calculation of OFP and SOAP

Different VOC species have different O_3 formation potentials, which are determined by reaction kinetics and associated mechanisms (Atkinson and Arey, 2003). The OFP has been used to evaluate the contribution of VOCs to O_3 formation in the troposphere, and subsequently, reactivity-based control approaches, which utilize the maximum incremental reactivity (MIR) values of each VOC species (Carter, 2010), i.e. Eq. (4).

$$OFP = \sum (c_i MIR_i) \tag{4}$$

where c_i is the measured concentration of VOC *i*; and MIR_{*i*} is the maximum incremental reactivity of VOC *i*.

Similar to the OFP, the SOAP can also be estimated according to the equation (5) (Derwent et al., 2010).

$$SOAP = \sum (c_i SOAP_i) \tag{5}$$

Thirty-nine VOC species with $SOAP_i$ values available in the literature were selected for the SOAP calculations in this study. The $SOAP_i$ represents the potential of a VOC species to form SOA relative to the same mass of toluene when it is added to the ambient atmosphere.

2.5. Box model for O_3 production rate calculations

The O_3 production rates were simulated using a box model with the master chemical mechanism (MCM 3.3.1) (Jenkin et al., 2017). The

diurnal curves of the O_3 concentration and production rate were calculated with the initial concentrations of propylene-equivalent VOC and NO_x , which were constrained by the hourly mean values of temperature, relative humidity, pressure, J_{NO2} , and concentrations of CO, SO₂, and HONO during our observations. The measured VOC concentrations were converted to equivalent concentrations of propene according to the MIR values (Zhang et al., 2020b). Isopleths (i.e., empirical kinetic modeling approach, EKMA curves) were drawn based on the daily maximum O_3 production rate. It should be noted that HONO can significantly promote O_3 formation, and its concentration was considered in the model simulations (Zhang et al., 2016).

3. Results and discussion

3.1. Overview of air pollution during observations

Fig. 2 shows the time series of air pollutant concentrations and the meteorological variables during the campaign. The temperature ranged from 21.4 to 34.4 °C, with a mean value (±standard deviation) of 27.4 \pm 2.8 °C, and the RH was 18.4%–93.0%, with a mean value of 57.3 \pm 19.1%. The mean wind speed was 1.7 ± 1.1 m s⁻¹ (ranged from 0.3 to 6.8 m s⁻¹), and the wind was predominantly from the southeast and southwest directions. The mean O3 concentration during the observation period was 44.8 \pm 27.2 ppbv, with a maximum of 119.1 ppbv, which is generally consistent with the mean O_3 concentration (49.1 \pm 29.1 ppbv) during 2014-2018 (Maji et al., 2020). The mean PM_{2.5} concentration was 17.4 \pm 21.4 μ g m⁻³, which was nearly within Grade II of the Chinese national ambient air quality standard (35 μ g m⁻³). Compared to 2014, in the suburban area, the concentration of PM_{2.5} decreased by 80.3% (Maji et al., 2020), reflecting the effectiveness of air pollution control policies in Beijing. When the wind speed is less than 2 m s $^{-1}$, local emissions followed by chemical transformation are usually the most important contributors to air pollution in Beijing (Wang et al., 2010). However, light PM pollution still occurred sometimes. O3 and PM_{2.5} pollution often occurred simultaneously (Fig. 2A and B) when the air masses originated from the south and wind speed was greater than 2 m s⁻¹ (Fig. 2F). From August 6 to 9 and from August 18 to 21, the pollutant concentrations (VOCs, $PM_{2.5}$, and O_3) increased substantially. Several O₃ pollution events occurred along with strong southwest wind after August 19. This suggests that regional transport also played important role in the accumulation of O₃ at higher wind speed during



Fig. 2. Time series of the concentrations of O₃, HONO, PM_{2.5}, OA, NOx, VOCs, J_{NO2} and the meteorological variables including temperature, relative humidity, wind speed (WS), wind direction (WD), and atmospheric pressure observed in August in Beijing. (In A, the red parts of the c_{O3} indicate the O₃ concentration exceeded 74.6 ppbv according to the second level of the Chinese national ambient air quality standard.)

these periods.

The concentration of VOCs was weakly correlated with those of OA $(R^2 = 0.30)$ and PM_{2.5} $(R^2 = 0.40)$, but was not correlated with the O₃ concentration $(R^2 = 0.05)$. This can be explained by the differences in the sources and/or formation mechanisms of OA and O₃. As an important component of PM_{2.5}, OA accounts for 20–90% of the mass of the PM_{2.5} in the troposphere (Jimenez et al., 2009). In our study, the OA in the PM_{1.0} contributed 46.7 \pm 20.5% to the PM_{2.5} mass concentrations, and SOA contributed 72 \pm 17% to the OA mass concentrations during our observations. This value is consistent with previous results, i.e., SOA accounted for 74% of the OA concentrations in Beijing in the summer of 2015 (Duan et al., 2020).

In addition to the contributions of secondary sources, primary emissions and other particle components also complicate the observed variations in PM2.5 concentration. The O3 concentration usually reaches a periodic peak in the early afternoon due to the active photochemical reactions under high temperature and solar radiation conditions (Lu et al., 2019). VOCs can break the dynamic equilibrium of NO-O₃-NO₂ and promote the conversion of NO to NO₂, finally leading to an accumulation of O₃. The O₃ concentration is nonlinearly dependent on the VOC concentration. As is shown in Fig. 2A and C, the VOC concentration was usually lower when the O₃ concentration was higher than the second level of the national air quality standard (the red parts in Fig. 2A). Thus, in addition to the variations in the atmospheric boundary layer (Garland and Derwent, 1979), O₃ is generated by the consumed VOCs and is sensitive to solar radiation variations (Lu et al., 2019). It should also be noted that in the early morning, both the concentrations of VOCs and NOx were sometimes relatively low, while the peak value of the O₃ concentration still frequently exceeded 74.6 ppbv. For example, the mean concentrations of NOx and VOCs were 21.8 \pm 5.1 and 10.6 \pm 1.4 ppbv, respectively, from 6:00 to 8:00 a.m. on August 1; while the 8 h-max O_3 concentration was 103 \pm 5.9 ppbv with a mean wind speed of 0.84 m s^{-1} . Hence, it is difficult to directly connect the observed precursor concentrations to the O₃ pollution when the wind speed is low because the photochemistry results in the production of O₃ and the consumption of the precursors.

3.2. Measured and consumed concentrations and PICs of VOCs

During the observations, 51 VOCs were detected and classified into four categories, including 26 alkanes, 9 alkenes, 1 alkyne, and 15 aromatics. Fig. 3A–C shows the time series of the measured and consumed concentrations and the PICs of the VOCs, respectively. Fig. 3D–F shows the corresponding diurnal profiles, and Fig. 3G–I shows the relative contributions of these four types of VOCs.

The measured VOC concentrations ranged from 2.2 to 33.6 ppbv, with a mean value of 11.2 ± 5.7 ppbv. Compared to the same period (45.4 ± 15.2 ppbv) in August 2015 (Li et al., 2016), the concentration of VOCs in Beijing were effectively reduced. Moreover, the consumed VOCs and PICs varied from 0 to 52.3 ppbv and from 2.8 to 63.8 ppbv, with mean values of 3.4 ± 7.6 ppbv and 14.6 ± 8.4 ppbv, respectively. The measured concentration of VOCs was significantly lower than that of the PICs (P < 0.05; Table S2). The ratios of PICs to the measured concentration should be paid to the degradation of VOCs when discussing the formation of secondary pollutants via VOC oxidation.

As is shown in Fig. 3D, the measured VOCs exhibit a peak in the morning and evening and a minimum at noon due to the evolution of the mixing layer height (MLH), VOC emissions from traffic, and photochemical reactions of VOCs (Sarkar et al., 2016; Zhang et al., 2020a). This is consistent with the typical diurnal curves of VOCs observed previously (Li et al., 2016). However, the concentrations of the consumed VOCs exhibited an obvious increase in the daytime (Fig. 3E), which is similar to the diurnal variations in the consumed VOCs reported by Shao et al. (2011). This indicates that solar radiation and temperature



Fig. 3. Time series, diurnal curves, and percentages of (A, D, and G) measured; (B, E, and H) consumed; and (C, F, and I) photochemical initial concentrations of VOCs.

may play important roles in accelerating photochemical oxidation of VOCs (Rubin et al., 2006). The diurnal profile of the PICs (Fig. 3F) shows the combined effect of the measured and consumed concentrations, i.e., the PICs start to increase in the early morning and decreased slightly throughout the day until evening.

As is shown in Fig. 3G and I, alkanes were the most abundant species for both the measured VOCs and PICs, accounting for 62.1% of VOCs and 49.2% of PICs, respectively. However, in Fig. 3H, the alkenes account for 80.5% of the total consumed VOCs, followed by aromatics (13.3%), alkanes (6.2%), and alkynes (0.02%). It is well known that alkanes are important VOCs produced by vehicle exhaust and account for a large portion of the VOCs in the urban atmosphere (Song et al., 2019; Liu et al., 2020b). Moreover, fuel combustion also contributes to the ambient alkanes in the Beijing-Tianjin-Hebei (BTH) area (Guo et al., 2017). The concentration of alkanes is relatively high in many Chinese cities due to high emissions and the relatively low reactivity of alkanes (Cai et al., 2010; Mozaffar et al., 2020). Thus, alkanes were only underestimated by $6.3 \pm 6.9\%$ when photochemical reactions were taken into consideration (Table S3), which is much lower than those for alkenes (43.8 \pm 30.2%) and aromatics (18.1 \pm 18.2%) because they are more reactive with OH radicals. Gao et al. (2018) reported that the PIC of alkenes and aromatics in winter in the urban areas of Beijing increased from 16.8% to 22.3% and from 8.7% to 9.9% of the observed VOCs, respectively, compared with the measured concentrations. In our study, the alkenes and aromatics had higher proportions in suburban areas (Fig. 3I), accounting for 31.3% and 18.8%, respectively. This is ascribed to the stronger oxidation capacity in summer (i.e., in this study) than in winter (i.e., the study of Gao et al., 2018). As is shown in Fig. S3A, isoprene, propylene, and butene are highly reactive species in the consumed VOCs. This illustrates that the VOCs with high reactivity

will react rapidly with OH radicals in suburban areas, resulting in apparent photochemical losses. Table S3 shows the details of the measured and consumed concentrations, the PICs and the average underestimated values for each species. Here, 12 species were underestimated by more than 30%. In particular, isoprene accounted for 67.9% of the consumed VOCs and 18.0% of the PICs. The PIC of isoprene was underestimated by 61.9%, which is slightly lower than that calculated by Shao et al. (2011) in the summer of 2008 (73%). Furthermore, as a critical biogenic VOC, isoprene has a high MIR (10.6 g g⁻¹) (Carter, 2010). Thus, the contribution of the biogenic VOCs to O₃ formation cannot be neglected in addition to anthropogenic VOC emissions (Song et al., 2019).

3.3. Measurement-based OFP (OFP_m) vs. consumption-based OFP (OFP_c) and PICs-based OFP (OFP_{PICs})

Fig. 4A, C, and E show the time series of the OFP_m, the OFP_c and the OFP_{PICs} with mean values of 57.8 \pm 26.3, 46.5 \pm 109.9 and 103.9 \pm 109.4 ppbv, respectively. It has been reported that the average OFP was about 135–164 ppbv near ground level in the BTH area (Geng et al., 2020), which is similar to our OFP_{PICs} value. Although alkanes accounted for 62.1% of the total measured VOCs (Fig. 3G), they accounted for only 21.4% of the OFP_m (Fig. 4B) due to the relatively small MIR values. In contrast, the alkenes and aromatics accounted for 36.6% and 41.9% of the OFP_m, respectively. The contribution of alkenes, aromatics, alkanes, and alkenes to OPF_{PICs} was 56.8%, 30.3%, 12.8%, and 0.1%, respectively. As is shown in Fig. 4D and F, alkenes dominated the O₃ formation calculated using the consumed VOCs and the PICs. In addition, biogenic emissions were a major contributor to O₃ production (Fig. S3B). Isoprene accounted for 74.8% of the OFP_{cc}, while its



Fig. 4. Time series and percentages of (A, B) OFP_m, (C, D) OFP_c, (E, F) OFP_{PICs}.

contribution to the OFP_{PICs} was 27.7%, which is comparable with the value (about 23%) reported by Xie et al. (2008) during the same period in downtown Beijing. In addition, several studies have found that propylene, trimethylbenzene, butene, and xylene produced by vehicle emissions, solvent use, and petrochemical plants are also important contributors to O₃ formation (Song et al., 2007; Liu et al., 2008).

Fig. 5 shows the relationship between the calculated OFPs and the observed O₃ concentrations. As is shown in Fig. 5A, the OFP_m was slightly negatively correlated to the observed O₃ concentration (OFP_m = 63.5–0.13 × c_{O3} , $R^2 = 0.32$), while the OFP_m from 6:00 a.m. to 8:00 a.m. was positively correlated with the 8-h maximum O₃ concentration $(OFP_m = 0.68 \times c_{O3} + 13.7, R^2 = 0.52, Fig. S4)$. However, the VOCs were continuously oxidized to form O_3 during transport, leading to a continuous increase in the O₃ concentration. Thus, a linear correlation $(OFP_c = 1.24 \times c_{O3}, R^2 = 0.93)$ was observed between the OFP_c and the O₃ concentration in Fig. 5B. A linear response of the OFP_{PICs} to the O₃ concentration was also observed (OFP_{PICs} = 41.5 + 1.40 \times c_{O3}, R² = 0.87, Fig. 5C). The slightly negative dependence of the OFP_m on the observed O₃ concentration (Fig. 5A) may be resulted from the fact that more VOCs were consumed, or less VOCs were left when the O3 level increased. These results imply that the OFP_{PICs} or OFP_c should more directly reflect the chemistry related to O3 formation. Thus, the O3 concentration can be predicted from the $\ensuremath{\mathsf{OFP}}_{\ensuremath{\mathsf{PICs}}}$ or $\ensuremath{\mathsf{OFP}}_{\ensuremath{\mathsf{c}}}.$ This can be explained by the formation mechanism of tropospheric O₃ in the presence of VOCs. The production of O₃ is governed by photolysis of NO₂,



Fig. 5. Relationships between the OFPs and the binned O_3 concentrations. A) for OFP_m; B) for OFP_c; and C) for OFP_{PICs}.

while NO_2 is formed by the oxidation of NO by RO_2 and HO_2 without consumption of O_3 . Both RO_2 and HO_2 should be positively correlated with the consumed VOCs rather than the measured VOCs. However, we

consider the individual VOCs in isolation when calculating the OFP. The actual photochemical O_3 formation results from complex photochemical reactions in the atmosphere involving many VOC species and NO_x . Thus, competitive reactions between different VOCs and OH radicals cannot be reflected when we calculate the OFP using the measured VOCs, but it is accounted for when we calculate the OFP using the consumed VOCs or the chemical transport model and box model. Although the correlation between consumption-based OFP and O_3 concentration was slightly better than that of PIC-based OFP (Fig. 5), it is likely that the PIC-based OFP is more suitable for estimating O_3 pollution, as it reflects the maximum potential of O_3 formation in the atmosphere.

Box models are a powerful tool for understanding the chemical regimes of O₃ formation (Ehlers et al., 2016). To further understand the different relationships between the measured VOCs and PICs on O3 formation, empirical kinetic modeling approach (EKMA) curves were derived based on box model simulations (Jenkin et al., 2017). As can be seen from Fig. 6, the O₃ production rate can be divided into three parts. The left-upper corner is a VOC-limited regime, while the right-lower corner is a NOx-limited regime. The area between the left-upper and right-lower corners is considered to be a transitional regime. In Fig. 6A, the measured VOCs and NO_x were used to assess the sensitivity of O₃ formation. The majority of the data resides in the VOC-limited zone, while a small amount of the data (in particular, for those datasets with O₃ concentrations of greater than 74.6 ppbv) resides in the transition regime. If we only consider the dataset measured in the early morning (Fig. S5), almost all of the data fall within the VOC-limited regime, with a VOCs/NOx ratio of 0.63 \pm 0.53 (Fig. S5). This is consistent with previous work which found that O3 formation is mainly VOC sensitive in Beijing (Li et al., 2020). When the PICs of the VOCs and NOx were taken into consideration, the sensitivity was significantly shifted toward the transitional and NO_x-limited regimes (Fig. 6B). As was discussed above, the consumed alkenes (especially isoprene) and aromatics mainly



Fig. 6. EKMA curves of the ozone generation rate in summer in Beijing. A) for the measured concentrations of VOCs and NOx; B) for the PICs of VOCs and NOx.

contributed to this large difference. Thus, the O_3 concentration is more sensitive to the PICs than to the measured VOCs. Our results suggest that it may be more reasonable to use the PICs-based values rather than the measured values to assess the sensitivity of O_3 production. However, more work is required to revisit the O_3 formation sensitivity in the future. In addition, more attention should be paid to the contribution of biogenic isoprene to O_3 formation in the summer in Beijing.

3.4. Measurement-based SOAP (SOAP_m) vs. consumption-based SOAP (SOAP_c) and PICs-based SOAP (SOAP_{PICs})

Fig. 7A shows the time series of the SOAP_m, which ranges from 1.9 to 22.5 μg m $^{-3},$ with a mean value of 8.4 \pm 4.1 μg m $^{-3}.$ It is similar to the SOAP value (7.7 μ g m⁻³) in the summer of 2016 in Beijing reported by Li et al. (2020). Aromatics were the most important contributor to the SOAP_m (98.5%), followed by alkanes (0.8%) and alkenes (0.8%) (Fig. 7B). Similar to $SOAP_m$, the contribution of the aromatics to the $SOAP_c$ (1.9 \pm 6.7 μg m $^{-3})$ was up to 92%, while the alkenes accounted for 7% (Fig. 7D). The aromatics, alkenes, and alkanes contribute 97.3%, 1.9%, and 0.8% to the SOAP_{PICs} (10.3 \pm 7.4 µg m⁻³), respectively (Fig. 7F). Aromatics were the main contributors to SOAP_{PICs}, which is similar to the value of 98.2% reported by Li et al. (2020). The contributions of the alkenes to the SOAP in the source region were also underestimated when based on the measured VOCs, by about 67.7% for the alkenes (Table S4). The top 10 PICs-based species accounted for 81.8% of the total SOAP_{PICs} (Fig. S3C). Toluene, benzene, and xylene were the important contributors to SOA formation (Kroll and Seinfeld, 2008), accounting for 60.4% of the total SOAP_{PICs}. In addition, the SOAP_c/SOAP_{PICs} ratios for styrene and isoprene were 84.2 and 88%, respectively, due to consumption during transport. Thus, the contribution of these species to SOA formation cannot be ignored during observations. Photochemical loss is more important for estimating OFP than SOAP. This can be explained by the higher reactivity of alkenes than aromatics with OH radicals.

It should also be noted that the lifetimes $(1/k_2c_{OH})$ of highly reactive VOCs, such as isoprene and propene, greatly depend on the OH concentration. The daytime (8:00-17:00) OH concentrations ranged from 1.1×10^6 molecules cm^{-3} to 1.3×10^7 molecules $cm^{-3},$ with a mean value of 4.3 \pm 3.1 \times 10⁶ molecules cm⁻³, based on the modelled J_{O1D} using a previously described method (Liu et al., 2020a; Liu et al., 2020c). The corresponding photochemical ages were 0.01–16.9 h (1.7 \pm 2.4 h). This value is comparable with previously reported photochemical ages (Shao et al., 2011; Gao et al., 2018) and indicates that the calculated PICs of highly reactive VOCs may be overestimated and should be taken as the upper limit. In addition, this method does not account for the transport of pollutants. Thus, while useful in a qualitatively sense, the method cannot quantitatively predict O₃ or SOA pollution. Although there are some uncertainties in estimating the photochemical age (Parrish et al., 2007), our results provide important information for evaluating the effects of photochemical processes on O3 and SOA formation. It should be noted that O₃ also readily reacts with alkenes, leading to additional consumption of alkenes. However, when calculating the consumption of VOCs by O₃, similar to OH exposure, the O₃ exposure $(=c_{O3} \cdot t)$ is required. The OH exposure is calculated from the changes in the VOC ratio using a pair of VOCs that are predominantly consumed by OH (Section 2.3). However, it is difficult to select an ideal pair of VOCs that are mainly consumed by O_3 . As a result, we did not consider the contribution of O₃ to alkene consumption in this study, and further only measured biogenic isoprene emissions. This will lead to an underestimation of the contribution of the photochemical loss of VOCs to OFP and SOAP and suggests that other biogenic VOCs need to be investigated in the future.

4. Conclusions

In this study, we conducted measurements of air pollutants in the



Fig. 7. Time series and percentages of (A, B) SOAP_m; (C, D) SOAP_c; and (E, F) SOAP_{PICs}.

southern suburbs of Beijing in August 2019. The OFP and SOAP were described based on the reactivity of VOCs. Of the detected 51 VOCs, the concentrations of alkenes and aromatics were underestimated by 43.8% and 18.1%, respectively, followed by alkanes (6.3%) and alkynes (1.4%). The alkenes and aromatics contributed 56.8% and 30.3% to the OFP_{PICs} respectively. The key species for O₃ formation were isoprene and propylene, and the main contributors to SOAP_{PICs} were aromatics, accounting for 97.3%. The main species for SOA formation were toluene and xylene. The PICs-based OFPs were correlated with the observed O3 concentrations, suggesting that the PICs-based OFPs can be used to predict O₃ concentrations. The O₃ formation shifted toward the transitional and NOx-limited regimes based on the PICs. Photochemical loss is more important for estimating OFP than SOAP because of the higher reactivity of alkenes than aromatics with OH radicals. In addition, it should be pointed out that only the isoprene from biogenic sources was measured, and it is necessary to assess the contribution of other biogenic VOCs to O₃ and SOA formation in the future.

Authorship attribution statement

Junlei Zhan: Methodology, Data curation, Writing - original draft. Zeming Feng: Methodology, Investigation, Data curation. Pengfei Liu: Methodology, Data curation, Writing - review & editing. Xiaowei He: Methodology, Investigation, Data curation. Zhouming He: Methodology, Investigation, Data curation. Tianzeng Chen: Methodology, Investigation, Data curation. Yafei Wang: Methodology, Investigation, Resources, Data curation, Writing - review & editing. Hong He: Writing - review & editing. Yujing Mu: Methodology, Writing - review & editing. Yongchun Liu: Conceptualization, Methodology, Data curation, Writing - review & editing.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Supplementary data

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